

INTRODUCTION

Severe and prolonged droughts between 1961 and 1988, combined with increased demands for freshwater supplies in the United States, have resulted in a critical need to assess the potential for development of ground- and surface-water supplies. Rapid industrial growth and urban expansion have caused existing freshwater supplies to be used at or near maximum capacity. Begin in 1978, the Regional Aquifer-System Analysis (RASA) Program of the U.S. Geological Survey (USGS) initiated a systematic effort to study a number of the Nation's most important aquifer systems, which, in aggregate, underlie much of the country and represent an important component of the Nation's total water supply. The broad objective for each of the 26 studies in the program is to assemble geologic, hydrologic, and geochronological information, to analyze and develop an understanding of the system, and to develop predictive capabilities that will contribute to the effective management of the system.

In 1988, as part of the RASA Program, the USGS began a 6-year study of the ground-water resources of parts of 11 States in the Eastern United States (Swain and others, 1991). The study was designated the Appalachian Valley and Piedmont Regional Aquifer-System Analysis (APRASA). The APRASA team investigated ground-water resources primarily in the unglaciated part of the Valley and the Blue Ridge, the New England, and the Piedmont Physiographic Provinces (fig. 1). For the purposes of this report, the small area in the New England Physiographic Province that is within the study area in New Jersey and Pennsylvania was considered part of the Piedmont Physiographic Province. The results of the APRASA are contained in about 50 reports and abstracts, including reports on simulation of ground-water flow in three water types, this atlas, and chapters in Professional Paper 1422. These chapters include the summary (Chapter A), descriptions of recharge rates and surface- and ground-water relations (Chapter B), hydrogeologic terranes in the Valley and Ridge Physiographic Province (Chapter C), and ground-water geochemistry (Chapter D).

The purposes of this atlas are to summarize the hydrogeology, to describe an analysis of maps and well records, and to present a classification and map of the hydrogeologic terranes of the Blue Ridge and Piedmont Physiographic Provinces within the APRASA study area. Hydrogeologic terranes are defined for this atlas as regionally mappable areas characterized by similar water-yielding properties of a grouping of selected rock types. The hydrogeologic terranes represent areas of distinct hydrologic character. The terranes are intended to help water users locate and develop adequate water supplies and to help hydrologists interpret the regional hydrogeology.

Previous investigations provide maps and descriptions of the geologic units, describe the local quantity and quality of ground water within these units, and establish the statistical methods for comparing the water-yielding properties of these units. State geologic maps show the distribution of geologic units at a scale of 1:500,000 for Alabama (Osborne and others, 1989), Georgia (Lawson and others, 1976), North Carolina (Brown and Parker, 1985), and Virginia (Calver and Hobbs, 1965). State maps show geologic units at a scale of 1:250,000 for Maryland (Clawson and others, 1968), New Jersey (Lewis and Kummel, 1992), Pennsylvania (Berg and others, 1980), South Carolina (Overstreet and Bell, 1965), Tennessee (Hardenman, 1966), and West Virginia (Cardwell and others, 1968). Quadrangle geologic maps show geologic units at a scale of 1:24,000 for parts of Delaware within the APRASA area (Woodruff and Thompson, 1972, 1975). Many efforts have been published describing the ground-water resources of a county, parts of a county, multi-county area, or river basins.

The statistical methods used in this atlas are based largely on those used by Helzel and Hirsch (1992) and by Knopman (1990, p. 79). In his analysis of well records in the USGS Ground-Water Site Inventory (GWSI) data base, Knopman (1990) noted factors that must be taken into account when assessing the water-yielding potential of the rocks in the Valley and Ridge, the Blue Ridge, and the Piedmont Physiographic Provinces in Pennsylvania. Readers are referred to Helzel and Hirsch (1992) and Knopman (1990) for details regarding statistical methods.

HYDROGEOLOGIC SETTING

The Blue Ridge and Piedmont Physiographic Provinces are underlain by metamorphic, igneous, and sedimentary rocks, many of which have undergone several periods of metamorphism and intrusive activity. Although 35 rock types have been identified and mapped in the Blue Ridge and 53 in the Piedmont, four rock types are particularly important in the hydrogeologic setting of the study area. The bedding and foliation are usually folded and tilted, can exhibit variable orientations, and commonly intersect one another in systematic geometric patterns. Although most rocks have strong directional fabric, igneous rocks are isotropic. In the Blue Ridge, the rocks are generally more massive and less foliated. The rocks were subjected to uplift during the Cenozoic Era, and subsequent weathering and erosion opened or widened existing fractures or created new ones by fracturing. Fault zones of different types, scales, and orientations are common; some fault zones are characterized by an abundance of fractures.

Regolith (soil, sediment, weathered rock) covers the consolidated rock throughout the study area, and bedrock typically is exposed only in areas of rugged topography or in stream channels where erosion has removed the surficial material. The regolith either was formed in place by the weathering of the underlying bedrock (residual) or was deposited after being transported from the place of weathering (alluvium, colluvium, alluvium). Much of the regolith in the study area is fine-grained residuum that retains the structural features of the parent bedrock (residual). Residual material is thin (less than 5 ft) on shale, phyllite, and carbonate rock, but can be thicker (more than 5 ft) on rock that contain substantial amounts of resistant quartzite, or well-sorted, cemented quartzite sandstone. Glacial till with a sandy or clayey matrix covers much of the uplands in New Jersey and eastern Pennsylvania and is less than 3 ft thick. In major valleys in the same area, alluvium commonly is more than 30 ft thick and is coarse-grained, well-sorted, and derived directly or indirectly from glacial deposits. Although of small areal extent in comparison to the entire study area, this alluvium produces some of the best yields of water to wells of any geologic unit in the study area. Within and adjacent to the Blue Ridge Physiographic Province, thin, stony colluvium overlies residual or bedrock on hill and mountain slopes that are created by resistant siliclastic rock. The colluvium commonly grades downward into dissected alluvial terraces containing cobble gravel and sandy loam. In several places (Frederick Valley, Mountain City Shady Valley, for example), colluvial-alluvial aprons cover cavernous, carbonate rock, which discharges to large springs and maintains unusually large base flow during dry weather (Hollyday and others, 1991).

The water-storage and transmission characteristics of the bedrock and the regolith and the hydraulic connection between the bedrock and the regolith determine the water-yielding potential of the aquifers in the metamorphic and igneous rocks in the two provinces (fig. 2). Wells that penetrate fractured rocks in areas mapped by thick, unconsolidated regolith generally have lower yields than wells that penetrate bedrock in the same area. Porosity of the regolith is about 35 to 55 percent near land surface (Stewart, 1962) but decreases with depth as the degree of weathering decreases. Porosity of the bedrock is about 10 to 20 percent (Held, 1964). Because the regolith is more porous than the bedrock, recharge is stored primarily in the regolith and is released slowly to underlying bedrock fractures. The transition zone between the saprolite and the unweathered bedrock is often more permeable than the saprolite. If the transition is thick, significant recharge can occur, and water can flow laterally in this zone. Laboratory analyses of samples of saprolite in Maryland and Georgia indicate a wide range of hydraulic conductivity—0.0013 to 15 ft/d (Nutter and Ott, 1969). The abundance of connected fractures in the rock directly affects the yield of wells in the Blue Ridge. Because fractures act as conduits for ground-water flow, well yields are greatest where wells intersect fractures that are large, numerous, or both. Because the number and size of fractures decrease with depth in the Blue Ridge and Piedmont rocks, few wells in these two provinces are drilled more than 500 feet deep.

Several basins and valleys that are filled with sedimentary rocks lie within the expanse of metamorphic and igneous rocks in the Blue Ridge and Piedmont Physiographic Provinces and constitute less than 10 percent of the study area. Carbonate rocks underlie the Catoctin Valley section of the Piedmont in southeastern Pennsylvania, the Frederick Valley in central Maryland, and small valleys and coves in the Blue Ridge Physiographic Province in Tennessee and in the Piedmont in Alabama (fig. 1). Although poorly represented in the study data base, the limestone, dolomite, and marble in these valleys produce as much as 1,800 gal/min to large-diameter municipal and industrial wells from conduits developed in the dense rock.

Siliclastic and volcanic rocks fill a series of 15 elongated, down-faulted basins that lie within the Piedmont Physiographic Province in a discontinuous belt extending about 600 mi from New York into South Carolina. These basins were formed in Tertiary and early Quaternary time during the incipient rifting of the continent. Concurrently, they were filled with thick (up to 20,000 ft) sequences of limestone, red shale, sandstone, and conglomerate that were intruded later by diabase sills and dikes. Structurally, most geologic units strike northeast and dip between 5 and 40 degrees toward the main border faults, commonly on the northwest or southeast margins of the basins. Ground water in the siliclastic rocks is stored and transmitted primarily through a complex network of interconnected openings formed along bedding-plane partings, joints, and faults. To a lesser extent, water is stored and moves through interstitial pores. The three Mesozoic basins (fig. 1) near the northeastern end of the discontinuous belt of 15 Mesozoic basins contain fractured shale, siltstone, and sandstone which yield as much as 2,200 gal/min to individual wells. Regionally, well yields tend to decrease substantially from north to south along the belt of 15 basins; this decrease is accompanied by a decrease from north to south overall grain size of the sediments that were deposited in the basins. Within the Mesozoic basins, water-yielding openings generally decrease in number and width of opening as depth below land surface increases.

Because fractures tend to decrease in number and size with depth, investigators have assumed that well yields decrease with depth. Data from a combination of domestic and non-domestic wells were sufficient in several local areas in the Piedmont Physiographic Province to investigate the relation between well yield and total depth of well as part of the APRASA study. The method was to compute the average yield of wells grouped according to their total depth into classes covering 50-ft depth intervals and to plot the average yield against the average depth of the interval (fig. 3). The curves for Howard and Montgomery Counties, Md., Greenville-Spartanburg, S.C., and counties west of Atlanta, Ga., show neither increase nor decrease in average yield with increase in average depth. On the other hand, the curves for counties around and near Winston-Salem, Stateville, and Charlotte, N.C., show a sharp increase in average yield with increase in average depth to depths of 400 to 600 ft, followed by a decrease in average yield with increase in average depth. Four out of seven areas may be concluded to have average well yields that are substantially greater for wells completed between 400 and 600 ft below land surface compared to wells in the same area completed between 100 and 200 ft below land surface.

SELECTION AND ANALYSIS OF RECORDS AND VARIABLES

Hydrogeologic terranes were classified separately within the two physiographic provinces by relating rock type to the yield of hydrogeologic terranes. The hydrogeologic terranes were then mapped by relating the rock type within each hydrogeologic terrane within each province back to the mapped geologic units with the corresponding rock type. The yields of non-domestic wells were obtained from the GWSI database of the USGS in each State. The analysis of maps and well records in order to classify and map hydrogeologic terranes involved several choices and methods. These included appropriate map scales, map projections, identification of the principal rock type in each geologic unit, and the use of all, or simply non-domestic wells, in statistical analysis. The selected choices and methods are described below.

GEOLOGIC MAPS

State geologic maps exist for the entire study area (fig. 1). The State maps, at scales of 1:500,000 and 1:250,000, were selected over maps of smaller or larger scale, or more recent vintage, because they depicted rock type in addition to rock age at a convenient scale. Excepting Delaware, film-positive maps of the geologic units were obtained from the State geologic maps. The maps were scanned and the resulting vector files were converted to line spatial data coverages, which were extensively edited to remove extraneous information and to match coverages for adjacent States (Mesko, 1993). State boundaries were added to each State coverage using a single APRASA wide scale. All coverages were reprojected from their original coordinate systems into the Albers coordinate system and joined to make a single geologic-unit coverage of the APRASA area. Maps at a scale of 1:24,000 (less than statewide) were used to cover the small area of Delaware within the study area.

A polygon coverage, which represents the area covered by each geologic unit, was created from the geologic-unit line coverage. Each polygon was labeled with the geologic-unit name, map symbol, rock type, and additional attributes. The rock type was determined from the explanation on the published geologic maps. Brief explanations were supplemented with descriptions of geologic units in the lexicons of geologic names. A polygon coverage of rock types was created by merging geologic-unit polygons that were labeled with the same rock type. This simplified coverage was then used to identify the rock type associated with each well in the GWSI database.

WELL RECORDS

The complexity of the geology of the Blue Ridge and Piedmont Physiographic Provinces causes large variation in the water-yielding properties of the rocks in comparison to the water-yielding properties of sediments in the structurally and lithologically simpler Coastal Plain Physiographic Province. The range in yield of water to wells completed in any particular rock type can span several orders of magnitude and overlap the ranges of well yields in other rock types. Any analysis of this variation requires a large amount of data to describe differences in the yield of water to wells within each rock type and to test for significant differences in yield among rock types.

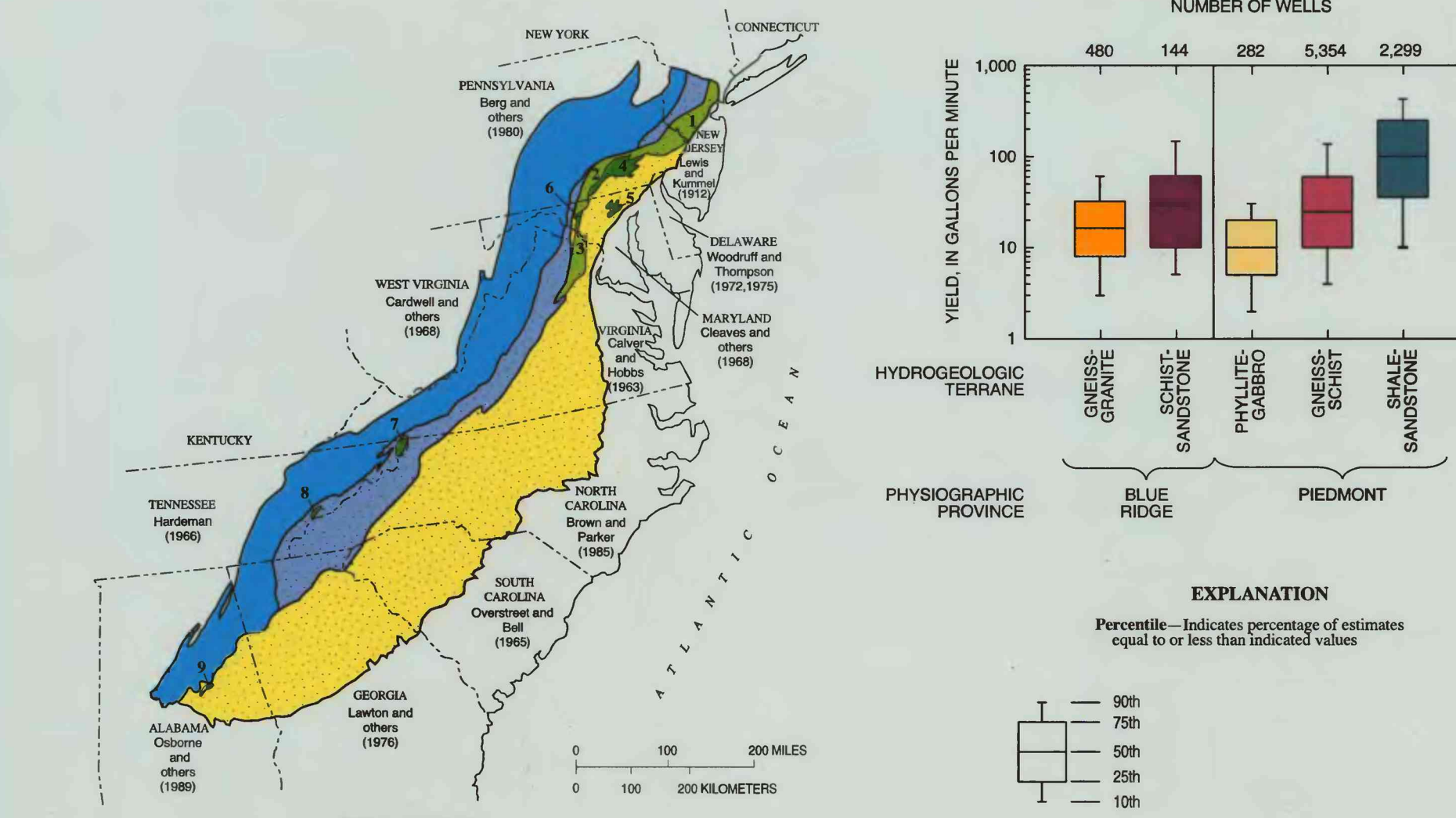


Figure 4. Variation of yield values for non-domestic wells grouped by five hydrogeologic terranes and two physiographic provinces.

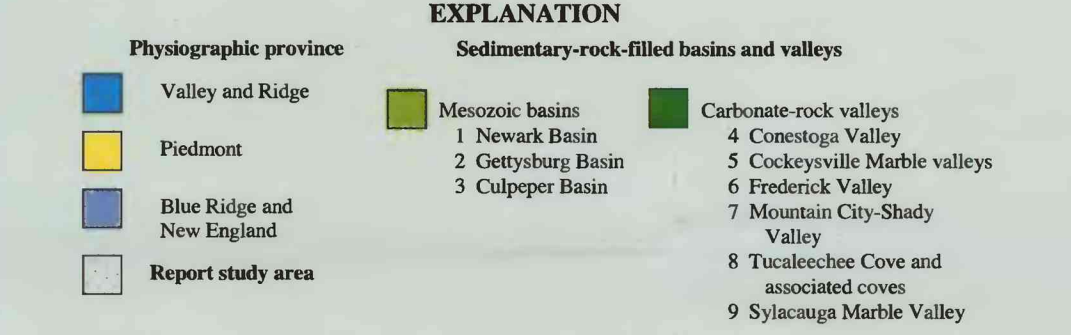


Figure 1. The Appalachian Valley and Piedmont Regional Aquifer-System Analysis study area. The physiographic provinces, report study area, sedimentary-rock-filled basins and valleys, and State-geologic-map references.

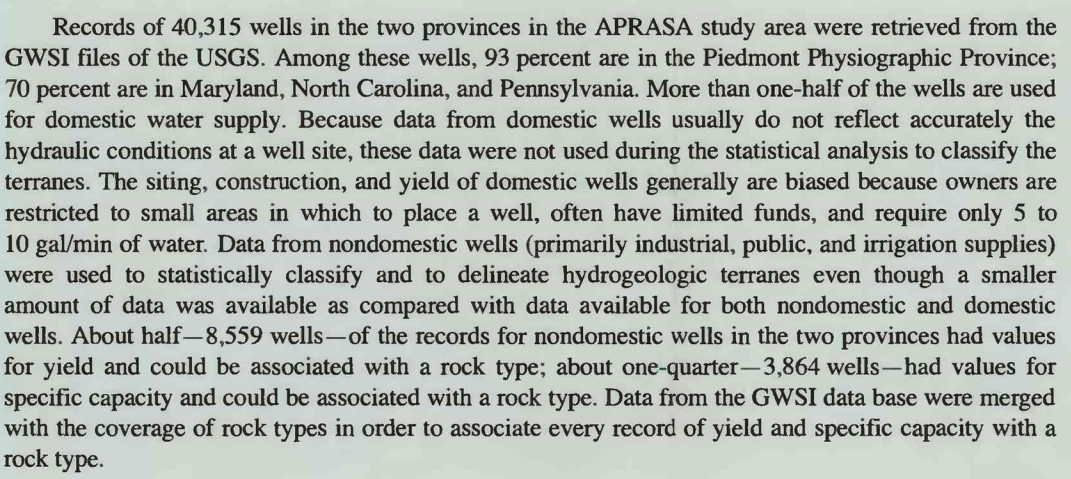


Figure 2. Generalized block diagram of the principal hydrogeologic components of metamorphic and igneous rocks in the Blue Ridge and Piedmont Physiographic Provinces (from Daniel and others, 1987, fig. 4).

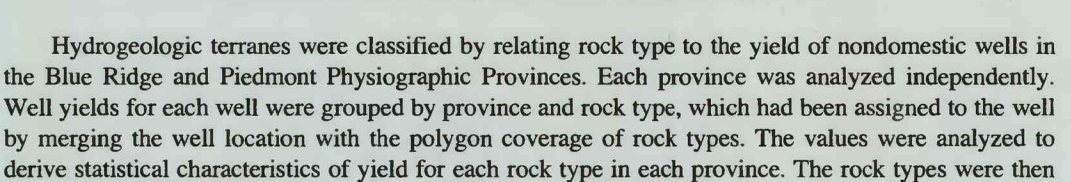


Figure 3. Relation between average well yield and average total depth of well for 50-foot depth intervals in seven areas of the Piedmont Physiographic Province.

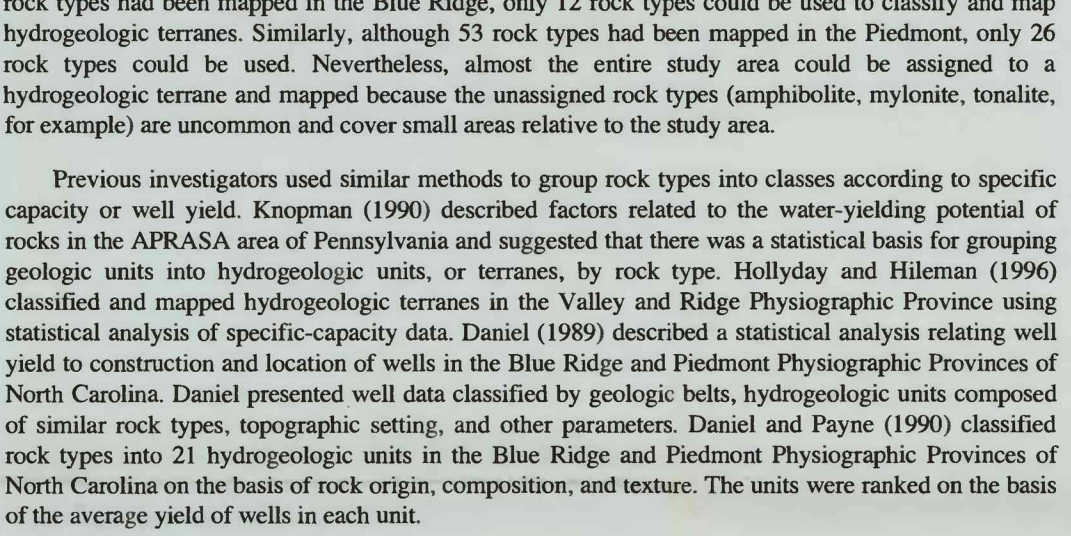


Figure 5. Variation of specific-capacity values for non-domestic wells grouped by five hydrogeologic terranes and two physiographic provinces.

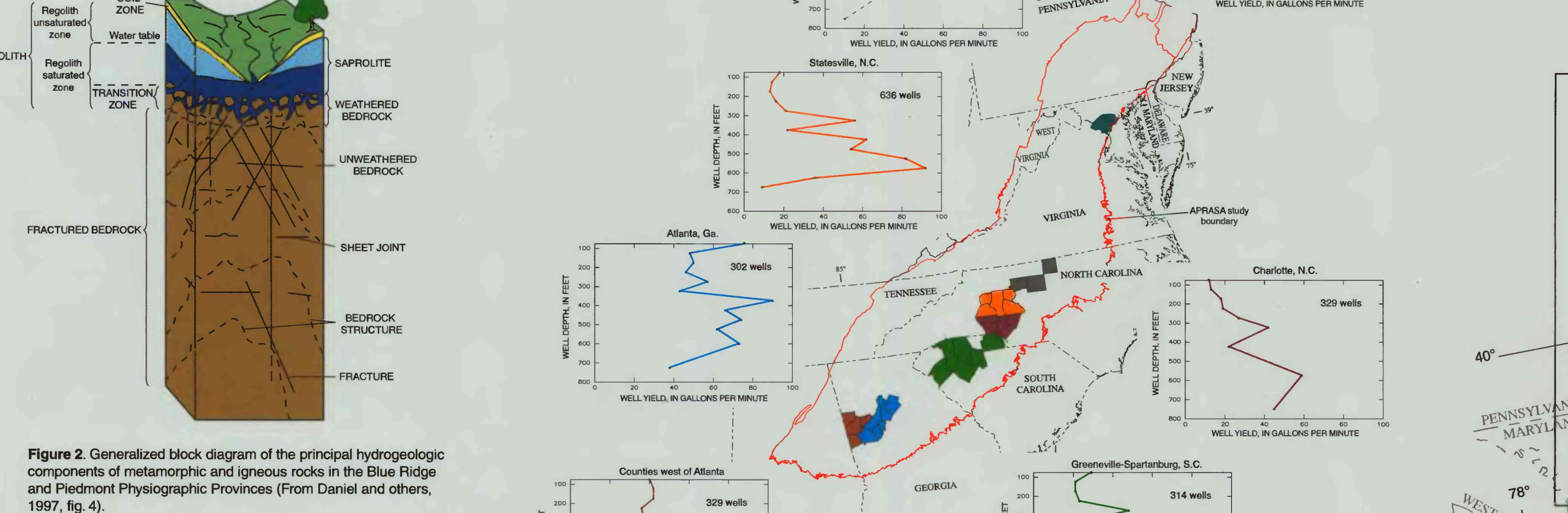


Figure 6. Location of the hydrogeologic terranes of the Blue Ridge and Piedmont Physiographic Provinces in the Eastern United States.

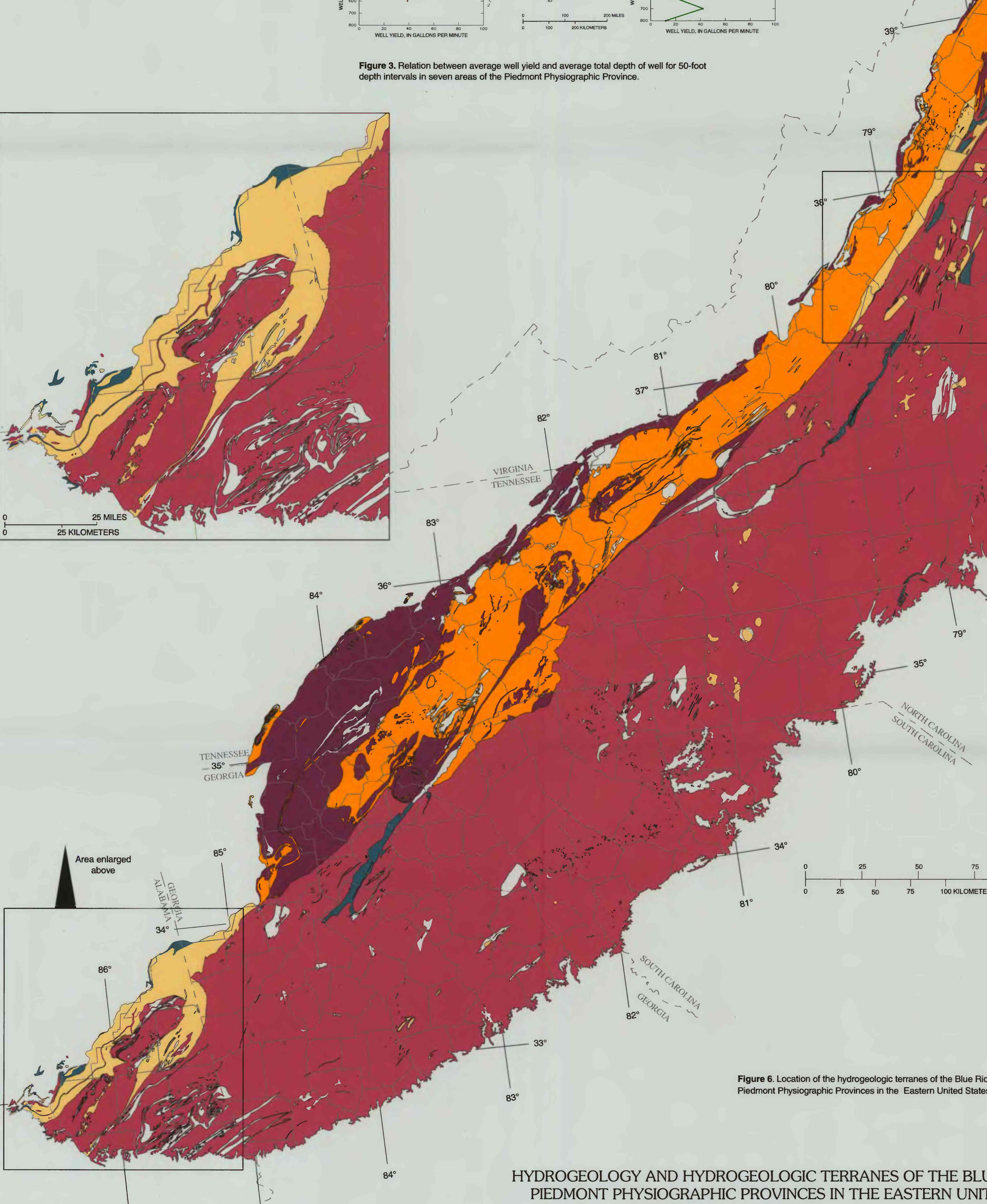


Figure 7. Hydrogeology and hydrogeologic terranes of the Blue Ridge and Piedmont Physiographic Provinces in the Eastern United States.

By T.O. Mesko, L.A. Swain, and E.F. Hollyday 1999

Table 2. Statistical characteristics of specific-capacity values of non-domestic wells in selected rock types in the hydrogeologic terranes of the Piedmont Physiographic Province

(Rock types are those that were determined by merging the well locations in the Ground-Water Site Inventory data base of the U.S. Geological Survey with the geologic information system polygon coverage of rock type. Specific capacity is in gallons per minute per foot of drawdown. The specific capacity is listed for P25, P50, P75, and the specific capacity at 25, 50, and 75 percent of the values are equal to or less than, respectively. Min. is the minimum specific capacity of record. Max. is the maximum specific capacity of record. Mean is the arithmetic mean, N, number of measurements. —, insufficient data to determine a value.)

Rock type	Min.	P25	P50	P75	Max.	Mean	N
Phyllite-gabbro hydrogeologic terrane							
Gabbro	0.01	0.09	0.28	1.1	20.0	10.0	11
Gneiss	0.02	0.07	0.20	0.51	2.5	0.9	11
Schist	0.01	0.03	0.14	0.37	1.0	0.39	82
Serpentine	0.01	0.06	0.22	0.71	7.5	0.97	12
Gneiss-schist hydrogeologic terrane							
Argillite	0.01	0.21	0.65	1.9	23	2.0	67
Conglomerate	0.00	0.35	0.93	2.5	130	5.3	88
Diorite	0.01	0.08	0.25	1.0	5.2	0.93	29
Gneiss	0.02	0.04	0.14	0.34	1.50	0.54	651
Granite	0.01	0.10	0.32	0.85	9.59	4.1	41
Quartzite	0.01	0.14	0.25	0.71	5.0	0.73	30
Schist	0.00	0.17	0.11	0.42	3.0	0.17	213
Shale-siltstone hydrogeologic terrane							
Basalt	0.02	0.17	0.14	0.28	3.0	0.43	43
Cryptocrystic	0.14	0.40	2.4	18	26	8.3	11
Granite	0.01	1.1	1.9	100	18	21	39
Sandstone	0.00	0.43	1.1	3.8	85	3.6	397
Siltstone	0.01	0.17	0.17	0.40	3.8	0.40	7
Siltstone	0.00	0.25	0.2	0.4	28	0.0	194

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38°

Area enlarged from left

37°